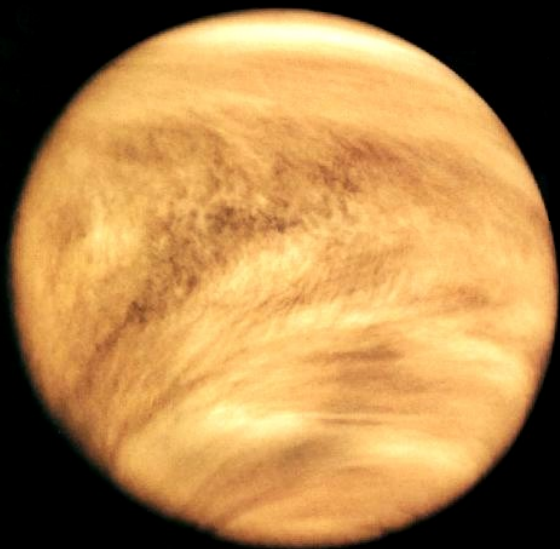


The mechanism of superrotation: Comparing Venus and Titan with General Circulation Models

S. Lebonnois

LMD, Paris



VENUS

$\langle T_s \rangle \sim 450^\circ\text{C}$

$\text{CO}_2 \sim 90\text{b}$

$\text{H}_2\text{O}/\text{CO}_2 \ll 1$

$\text{N}_2 \sim 3\text{b}$

Sun distance = 0.72 AU

$M = 0.81 M_{\text{Earth}}$

$\rho = 5.25$

obliquity = 177.4°

rotation = (-) 243 d

revolution = 224.7 d



EARTH

$\langle T_s \rangle \sim 15^\circ\text{C}$

$\text{CO}_2 \sim 0.3\text{ mb}$

$\text{O}_2 \sim 0.2\text{ b}$

$\text{N}_2 \sim 0.8\text{ b}$

1 AU

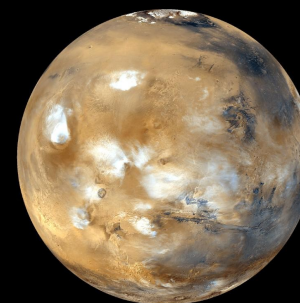
1

5.52

23.5°

23 h 56 m

365.25 d



MARS

$\langle T_s \rangle < -50^\circ\text{C}$

$\text{CO}_2 = 0.006\text{ b}$

$\text{N}_2 = 0.0002\text{ b}$

1.52 AU

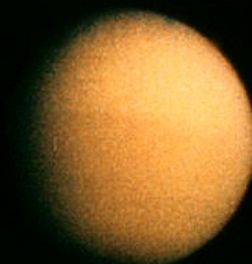
$0.11 M_{\text{Earth}}$

3.95

25.2°

24 h 37 m

687 d



TITAN

$\langle T_s \rangle \sim 95\text{ K}$

$\text{CH}_4 \sim 0.06\text{ b}$

$\text{N}_2 = 1.5\text{ b}$

9.5 UA

$0.023 M_{\text{Earth}}$

1.88

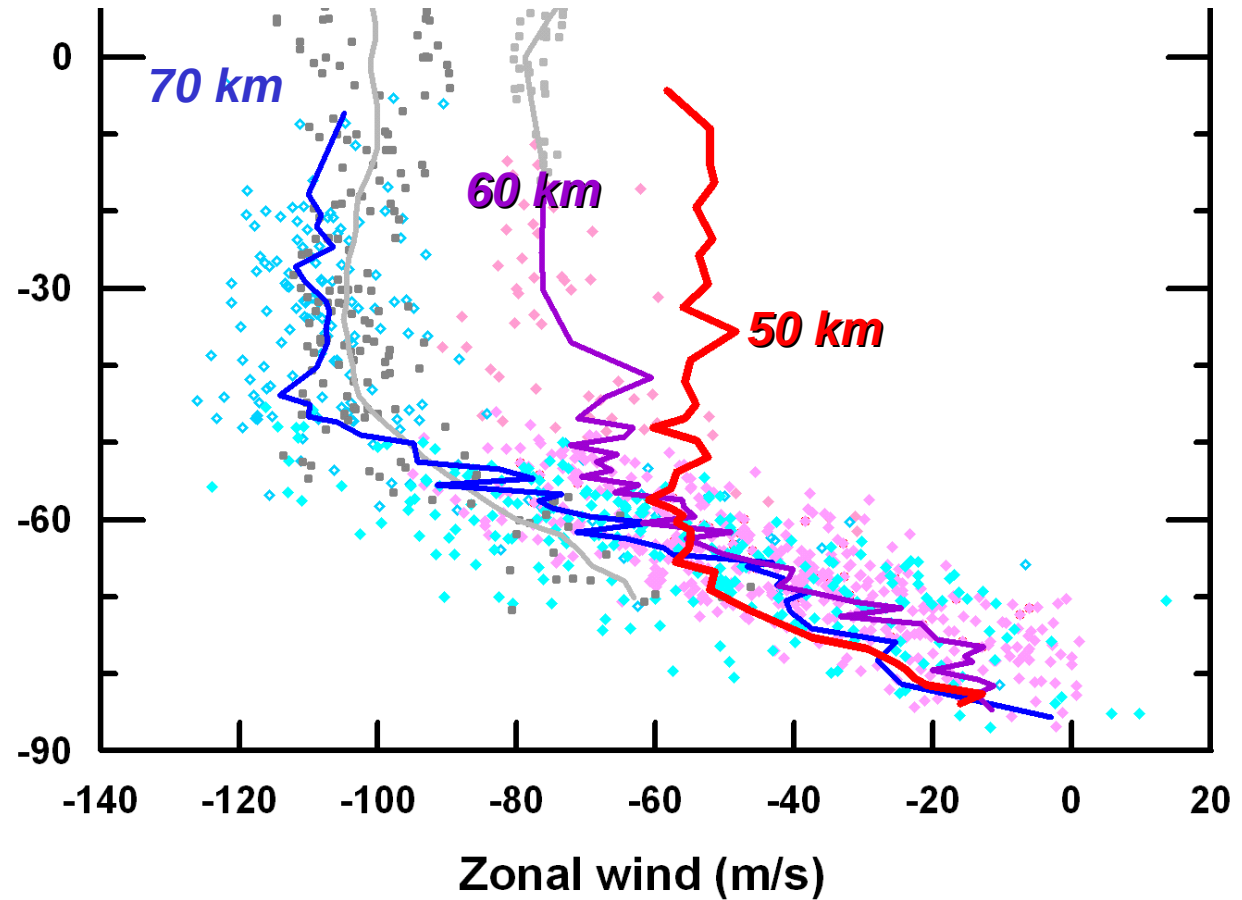
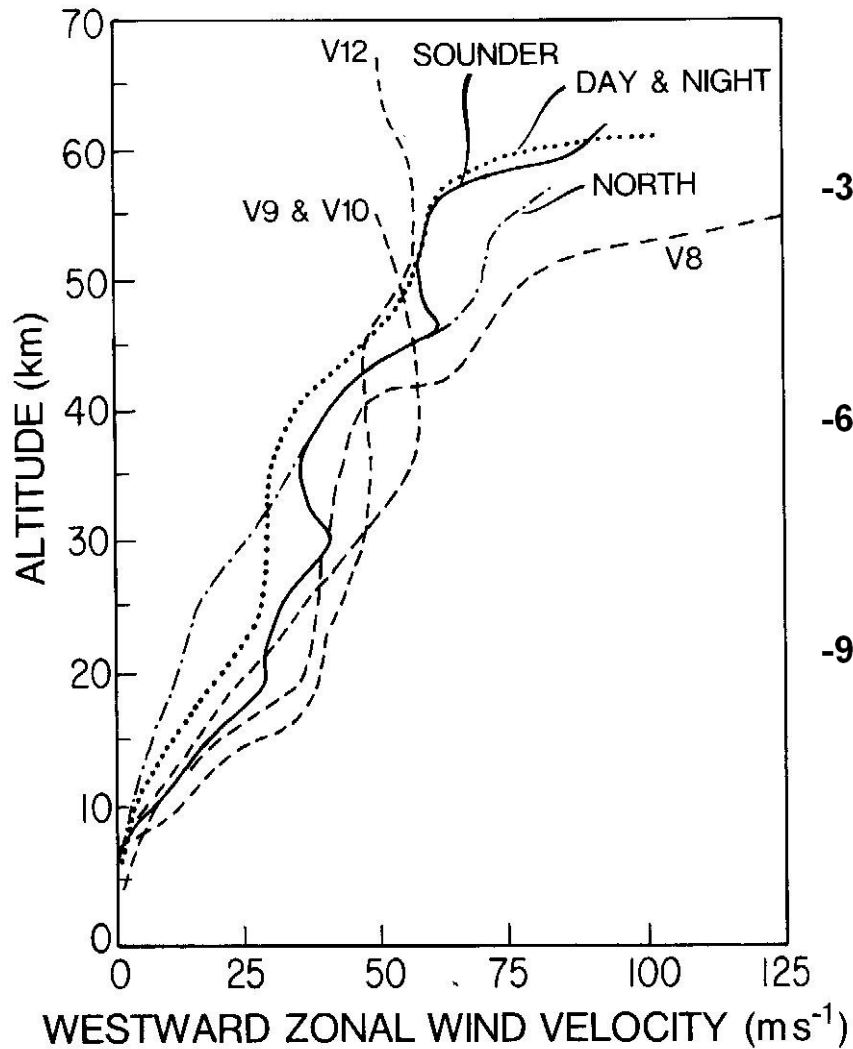
26.7°

15.94 d

$\sim 30\text{ years}$

Observations of superrotation

Pioneer Venus and Venera probes

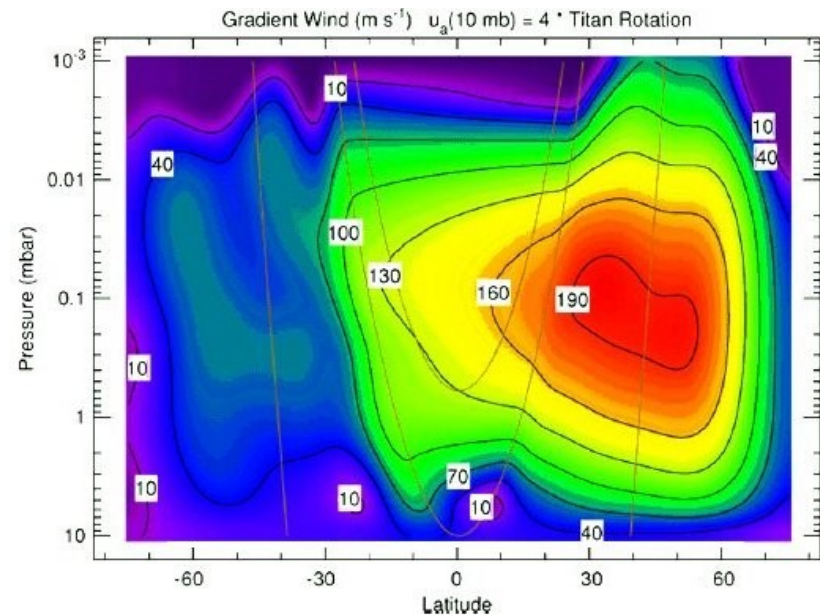
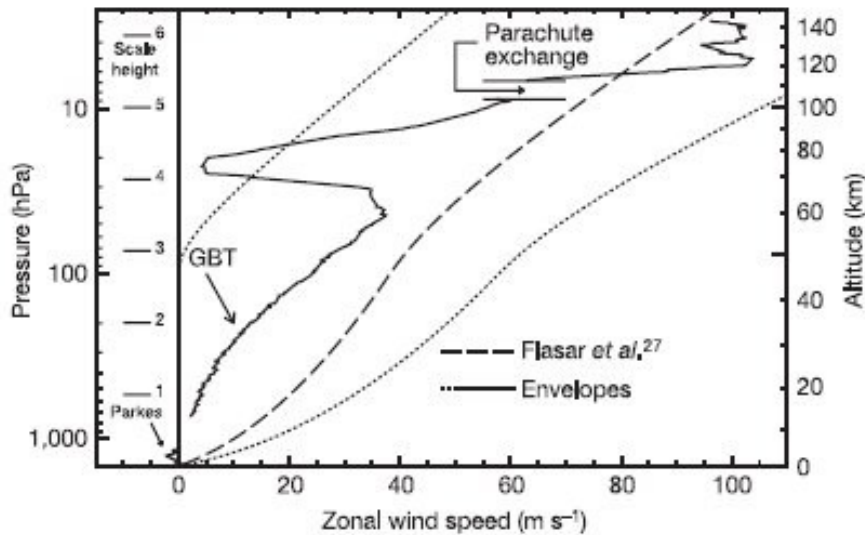
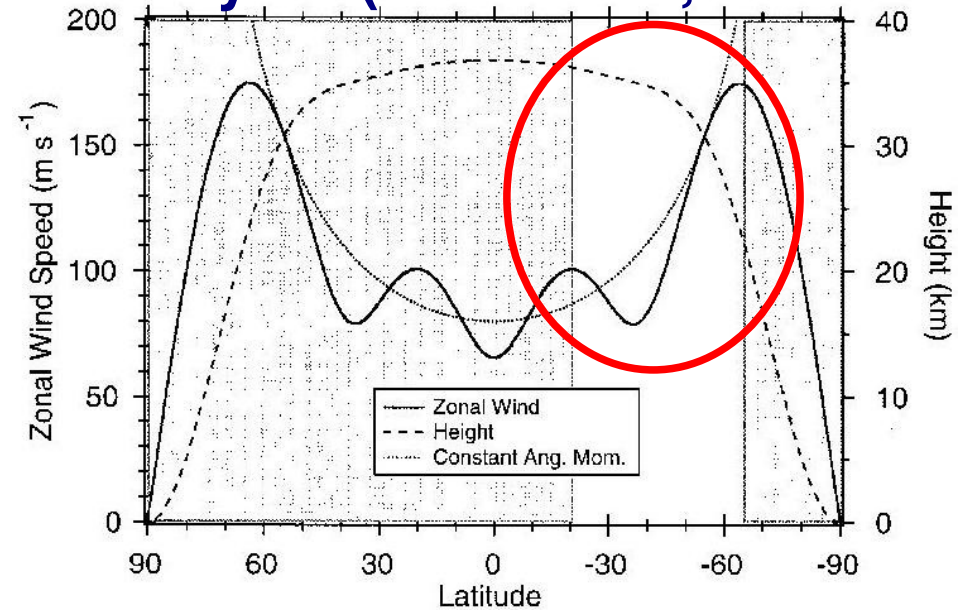


Venus Express/VIRTIS cloud tracking

Observations of superrotation

Stellar occultation analysis (0.25 mbar, $L_s \sim 128^\circ$)

Huygens/DWE vertical profile at $10^\circ S$ ($L_s \sim 300^\circ$)

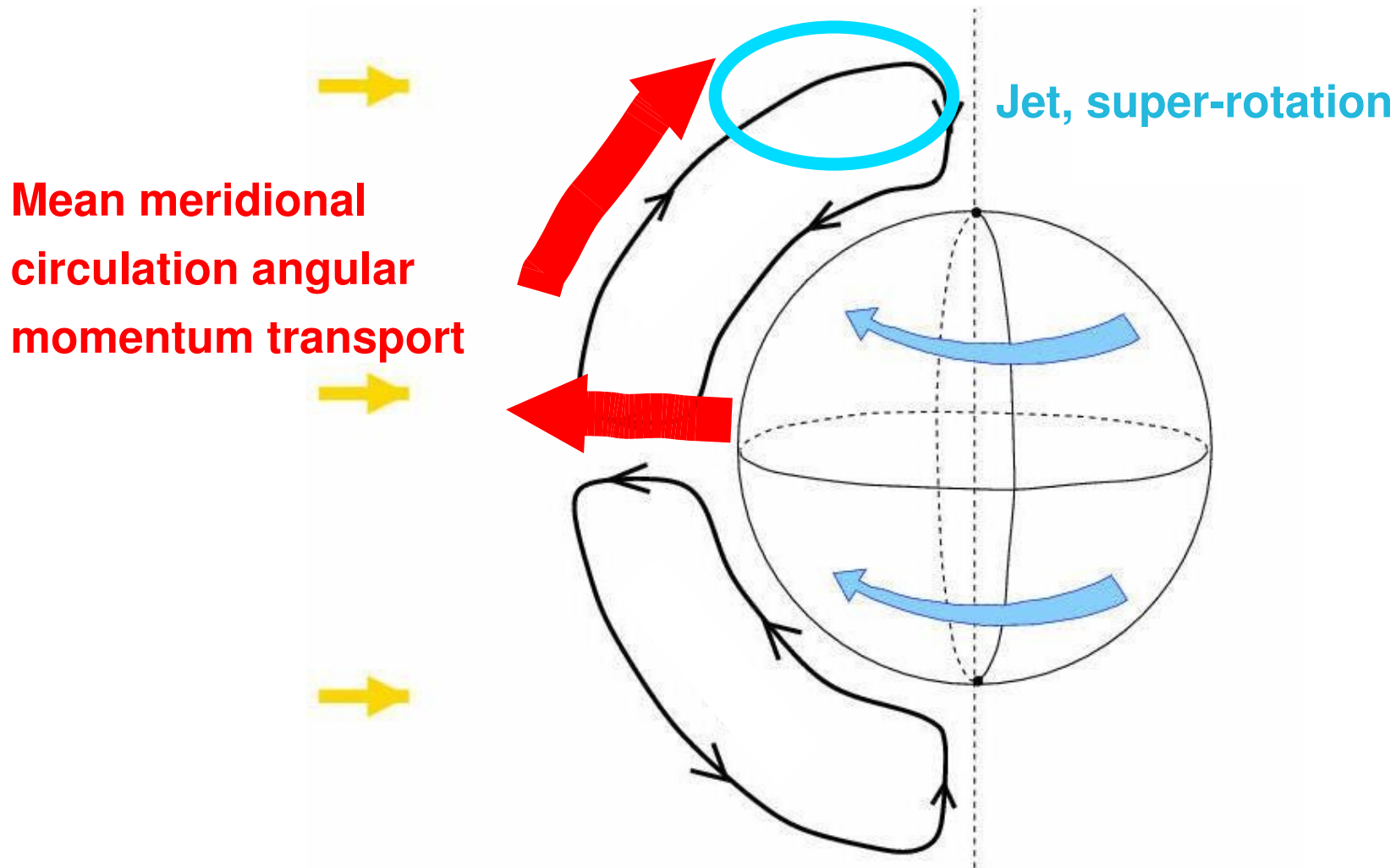


Cassini/CIRS thermal winds retrieval ($L_s \sim 300^\circ$)

The superrotation question

Slow rotation:

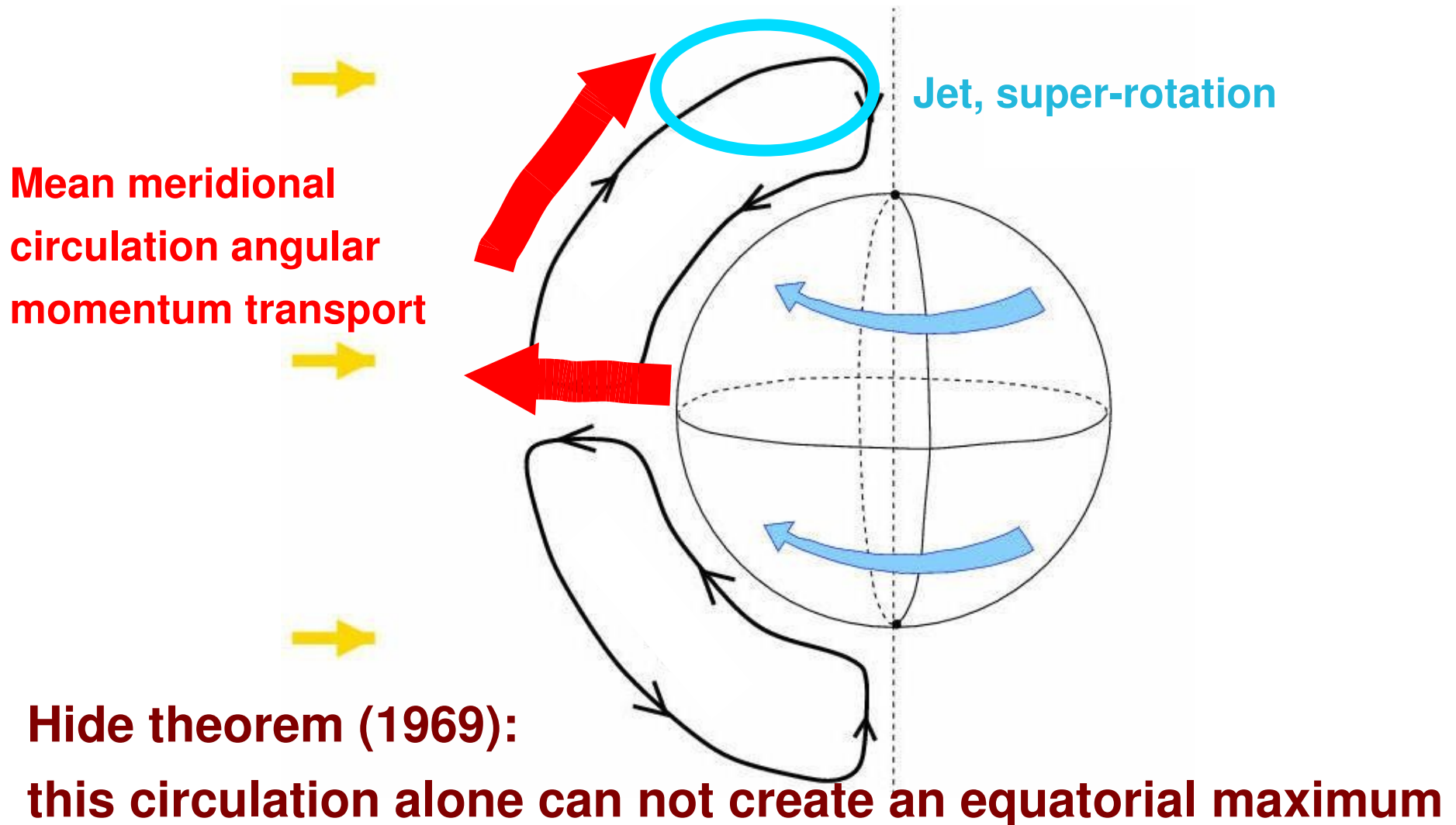
extension of Hadley cells from equator to the poles



The superrotation question

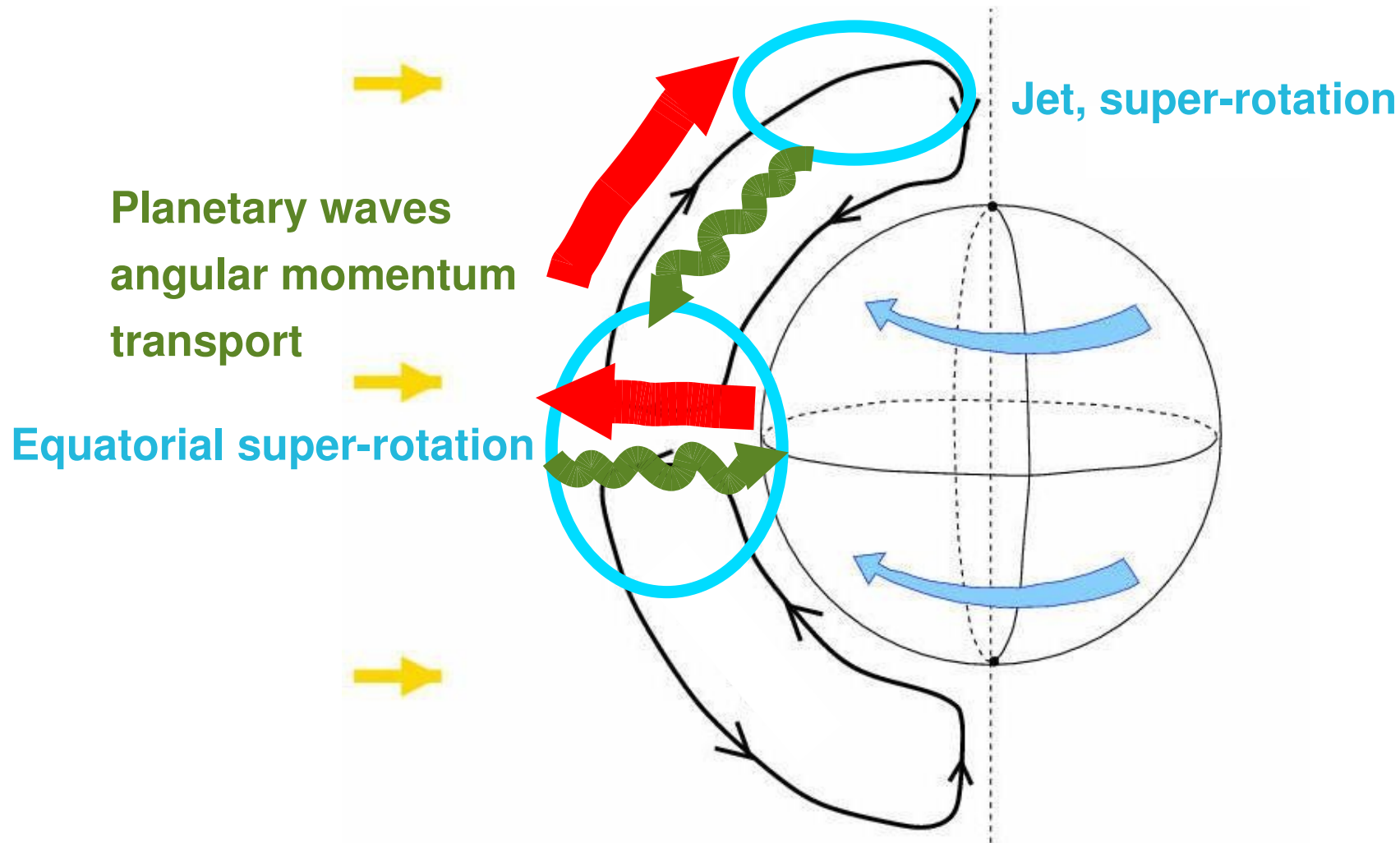
Slow rotation:

extension of Hadley cells from equator to the poles



The superrotation question

Superrotation at the equator:
need for non-axisymmetric planetary waves



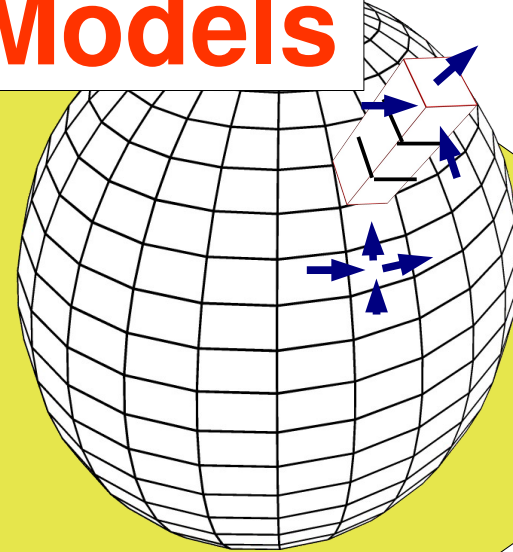
The superrotation question

- Gierasch 1975; Rossow and Williams 1979:
GRW mechanism: **unstable high-latitude jets**
Horizontal transport by **waves** from poles to equator
- Newman and Leovy 1992, Takagi and Matsuda 2007:
Possible role of **thermal tides**
Vertical transport of angular momentum
- Leovy 1973, Hou and Farrel 1987:
Possible role of **gravity waves**
Vertical transport of angular momentum

General Circulation Models

Dynamical core (3D or 2D)

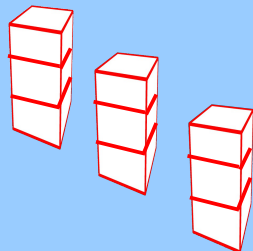
- Primitive equations of meteorology
- Used for weather forecast and climate
- Depends on a few parameters only (gravity, gaz molecular mass, planetary radius, R/C_p)
- Finite differences or spherical harmonics



$$U^*, v^*, T^*, P_s^* \xleftarrow[\Delta t \sim 5 \text{ min}]{\delta_t u^*, \delta_t v^*, \delta_t T^*, \delta_t P_s^*}$$

$$U, v, T, P_s \xrightarrow[\Delta t \sim 30 \text{ min}]{\delta_t u, \delta_t v, \delta_t T, \delta_t P_s}$$

Set of physical parameterizations specific of the planet



- radiation : main change
- subgrid scale processes
- specific processes (condensation, clouds, vegetation, ...)
- surface scheme

LMDZ versions

Earth
Mars
Titan
Venus
Idealized

First LMD Titan GCM

Hourdin et al. 1995

- Three-dimensional
- Fixed homogeneous haze and composition
- Surface to ~250 km

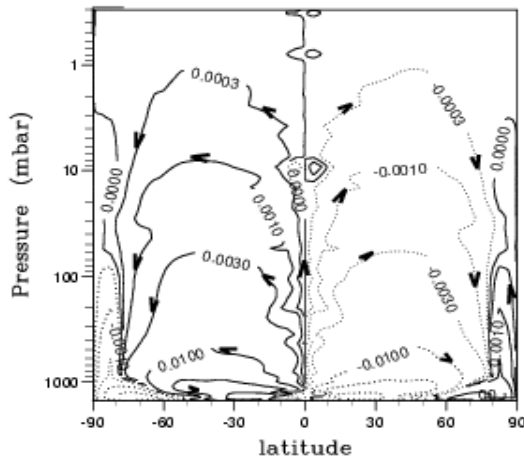
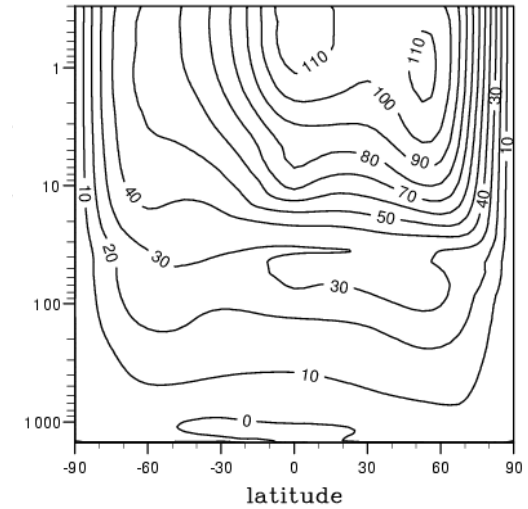
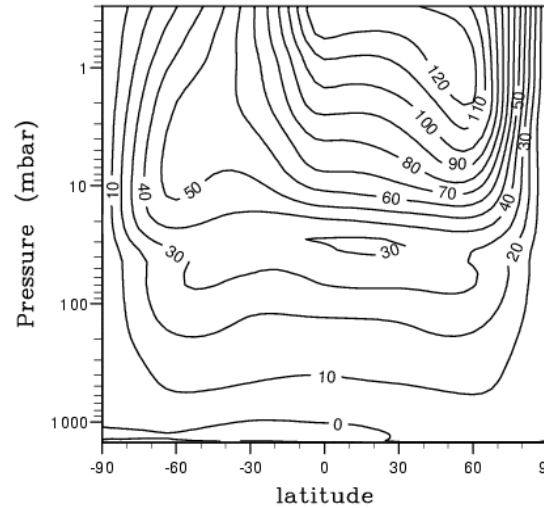
Southern summer solstice



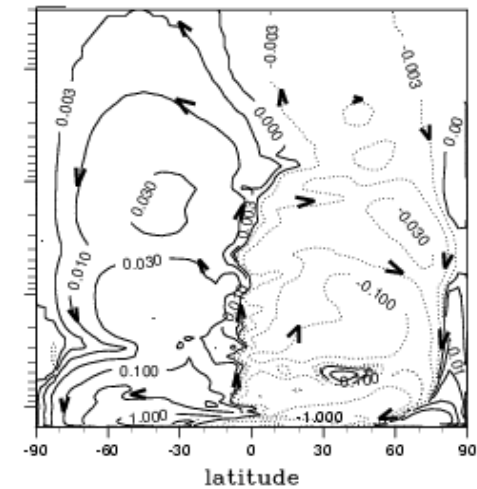
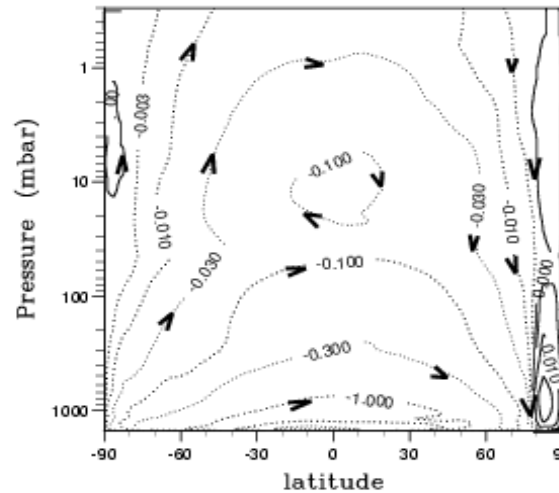
Northern spring equinox



Zonal wind (m/s)



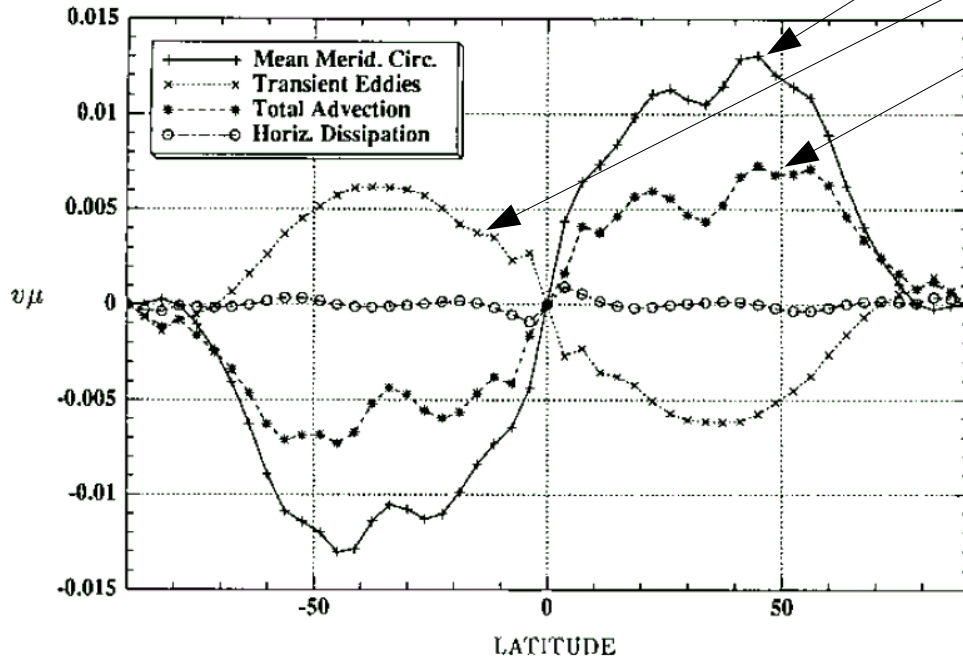
Stream function of the Mean
Meridional Circulation (Mt / s)



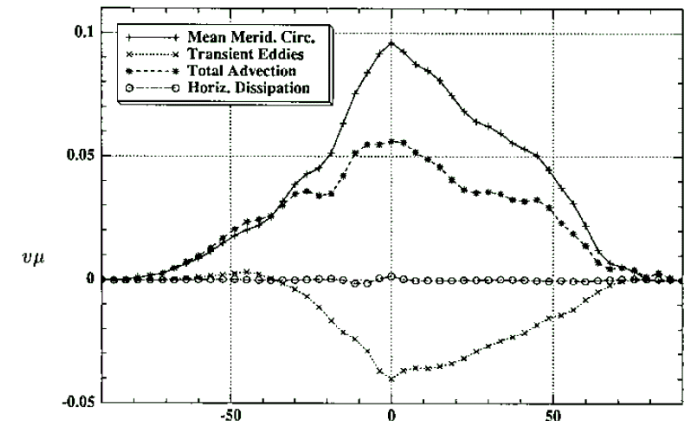
First LMD Titan GCM

Hourdin et al. 1995
annual mean

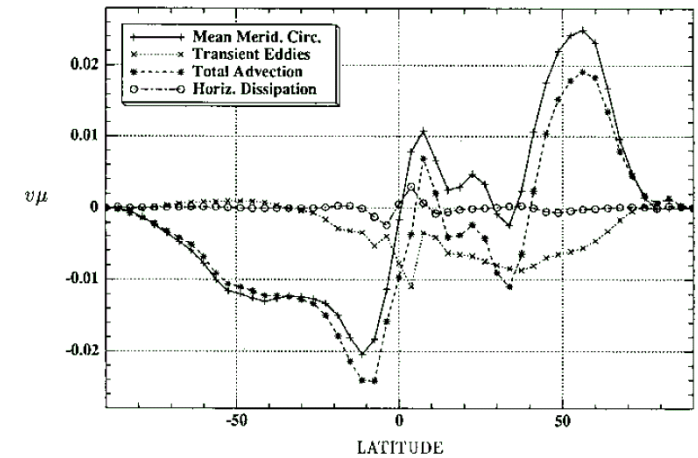
Mean meridional circulation
Transients
Total advection



N winter solstice



⇒ GRW mechanism



N spring equinox

Coupling with haze and chemistry

2D Climate Model

- **Barotropic waves have to be parameterized**
- Coupled haze and composition
- Surface to ~500 km

Important step :

Development of a parametrization of latitudinal mixing by waves.

PhD work of David Luz

1. Study of the **mixing properties of barotropic planetary waves** in Titan stratosphere (Luz et Hourdin 2003).

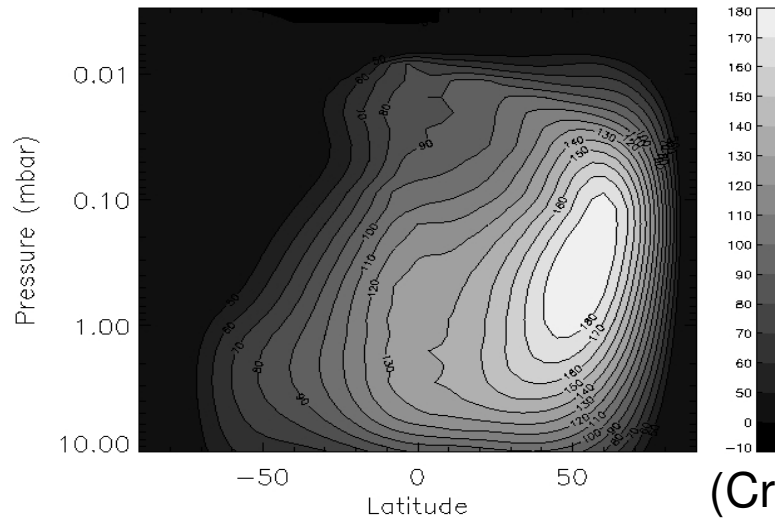
2. Development of a **parameterization** (Luz et al., 2003).

Done with a **2D longitude-latitude "shallow water"** model.

Coupling with haze and chemistry

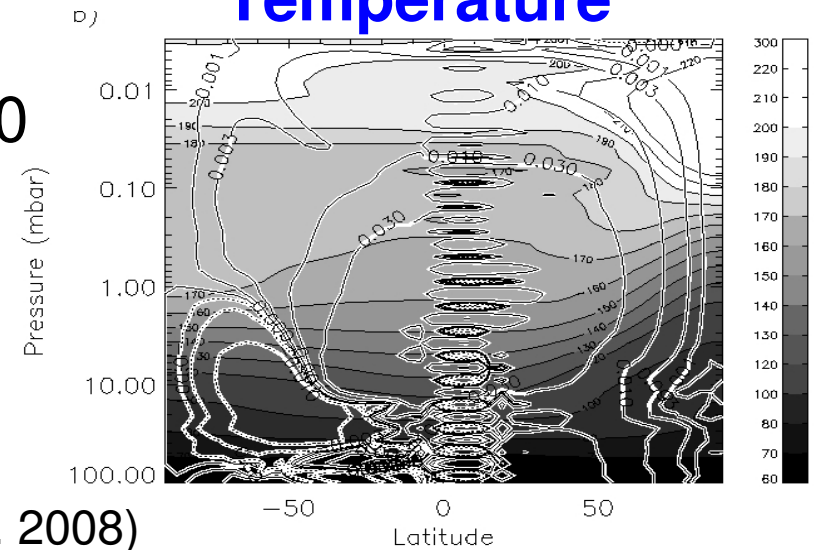
2D Climate Model

Zonal wind



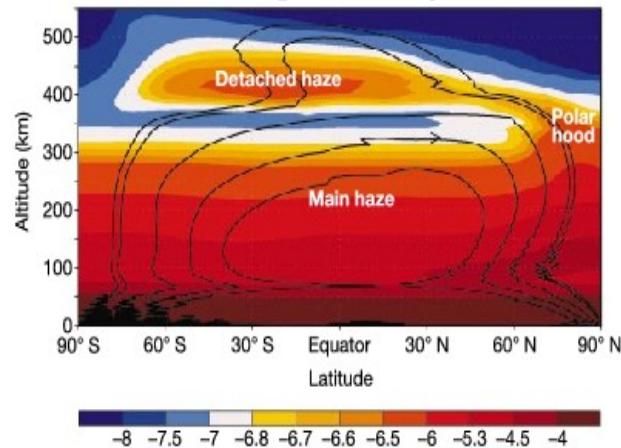
LS=300

Temperature

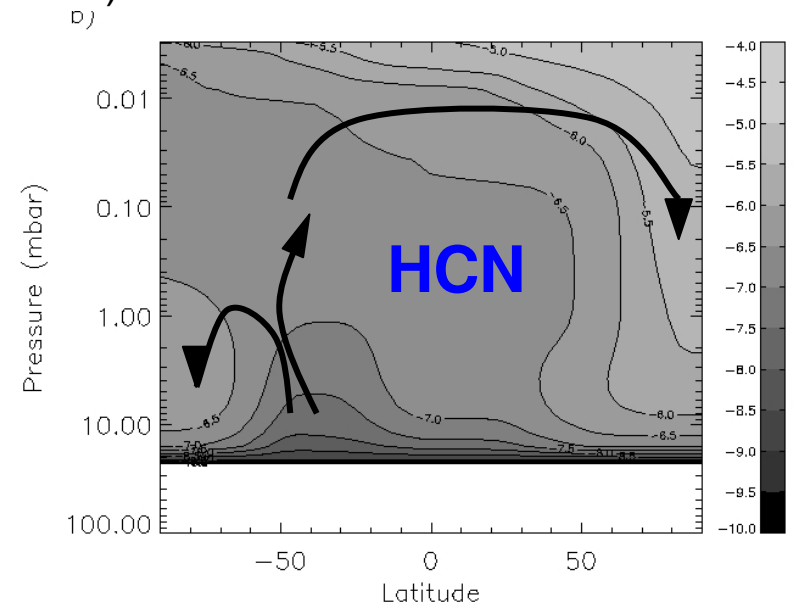


(Crespin et al. 2008)

Haze opacity



(Rannou et al. 2002)

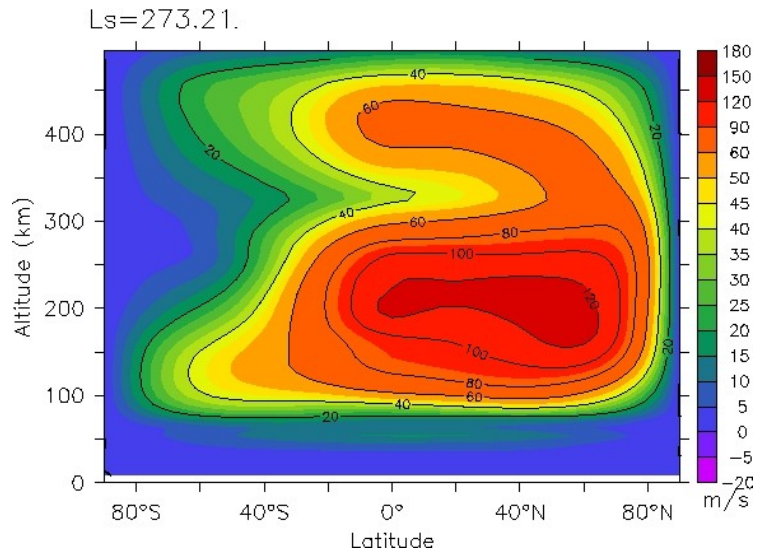


Back to 3D GCM

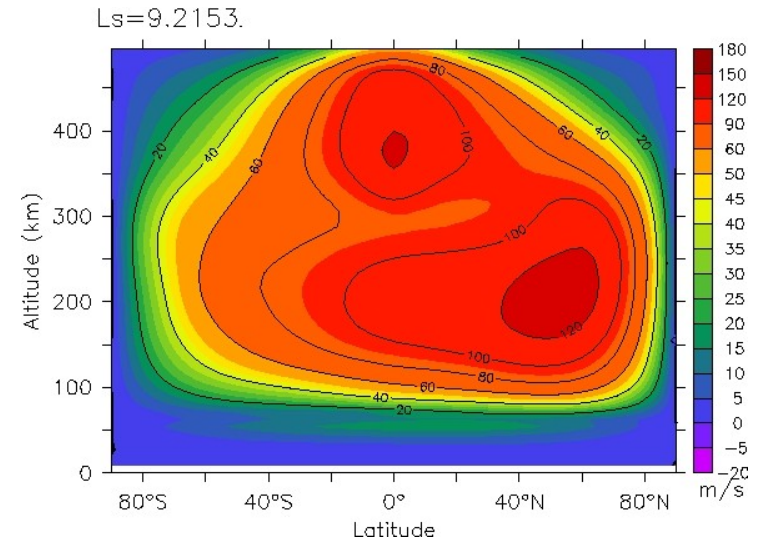
- 48x32x55 (0~500 km)
- Haze microphysics coupled
- No clouds microphysics coupled
- No photochemistry coupled
- Diurnal cycle
- Starting from 2D simulation
- 5 Titan years run
- Structure obtained similar to old Hourdin et al. (1995) 3D simulations

Back to 3D GCM

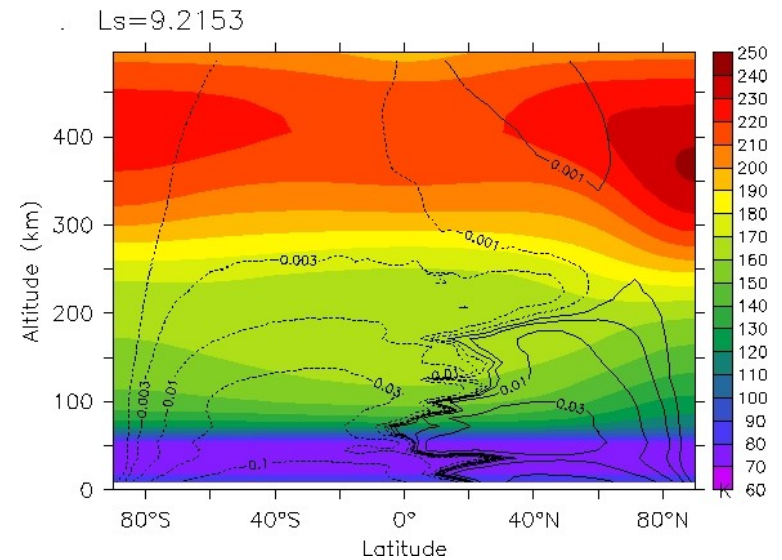
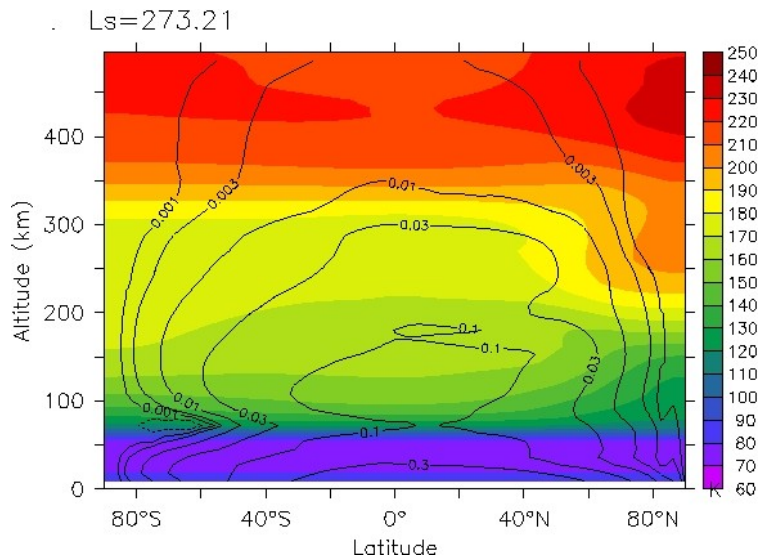
Mean zonal wind



N winter solstice

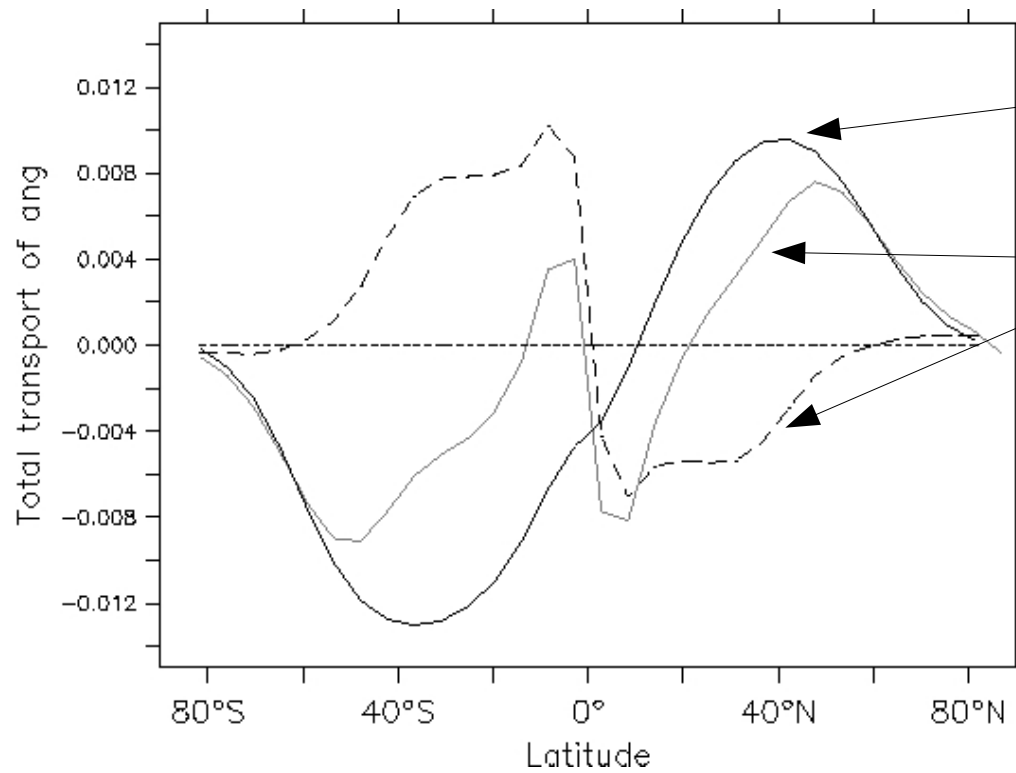


N spring equinox



Mean temperature and stream function

Momentum transport

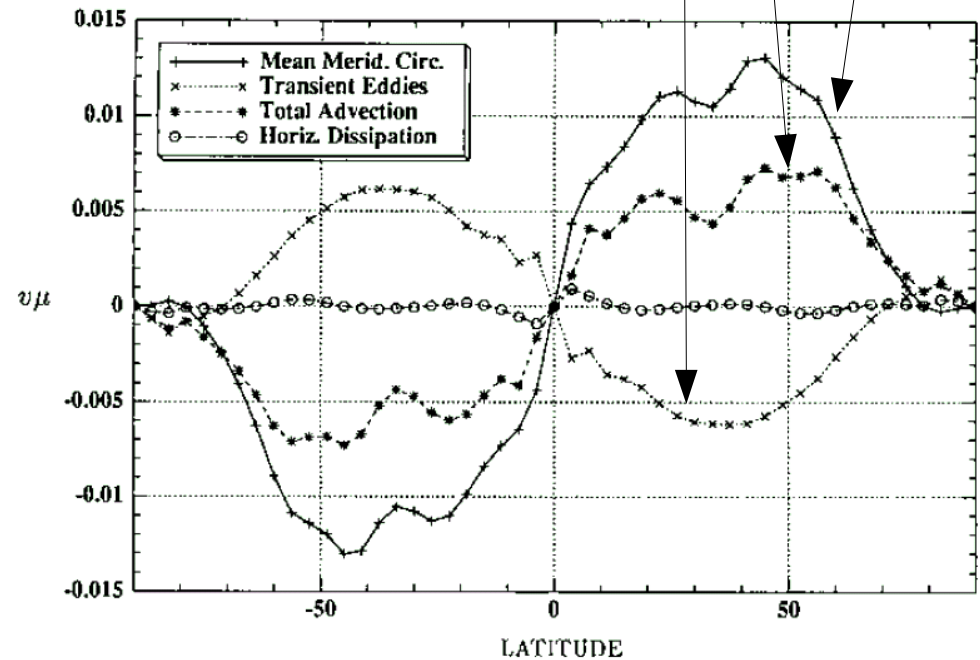


3D, annual mean

⇒ GRW mechanism

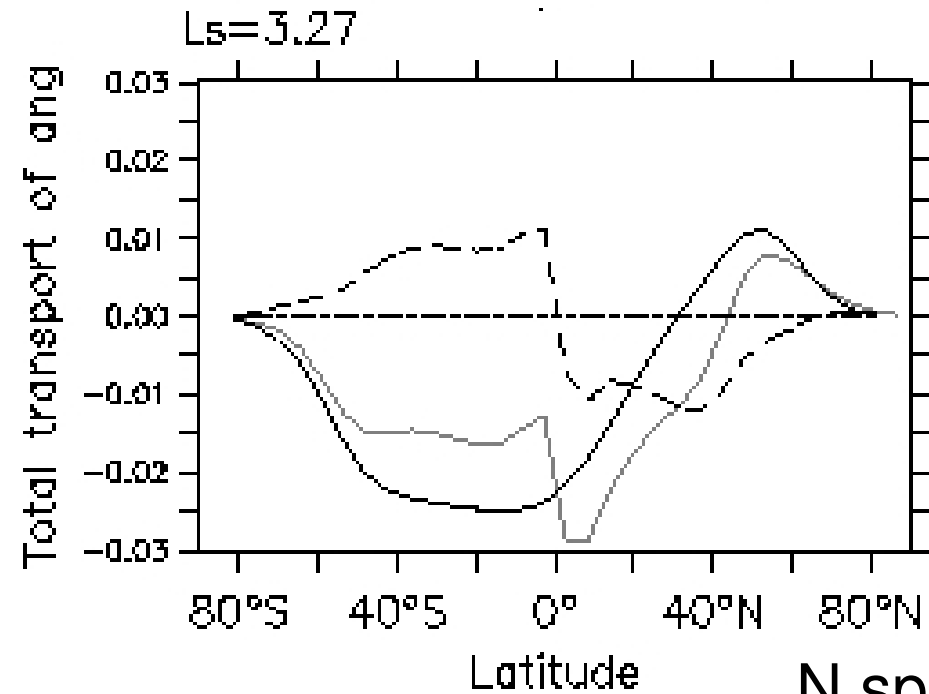
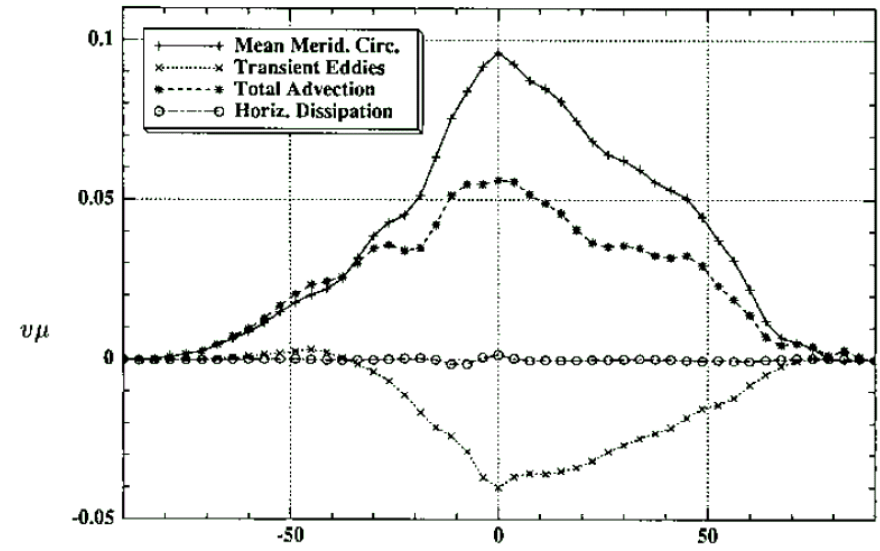
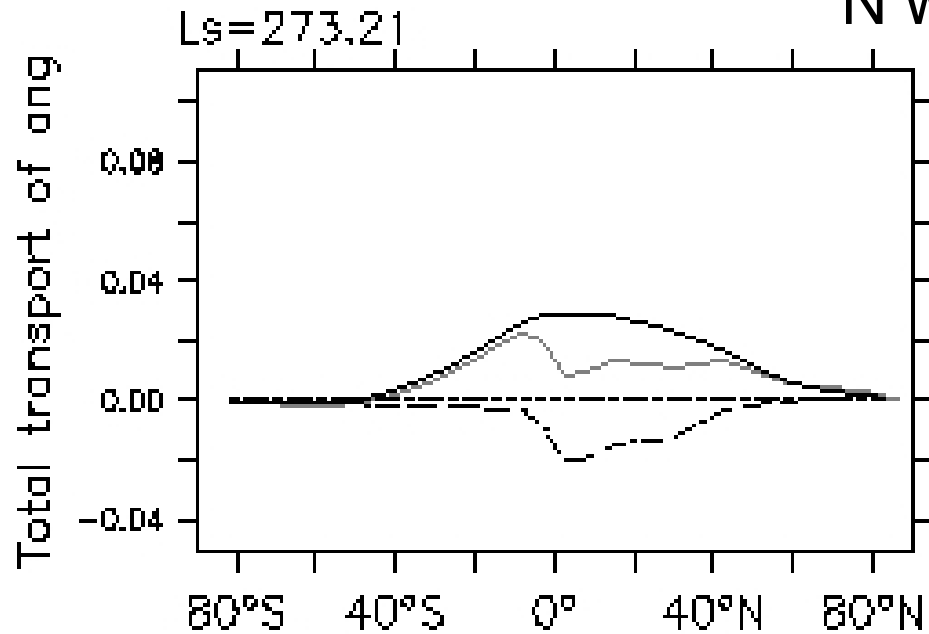
Mean meridional circulation
Transients
Total advection

Hourdin et al. 1995,
annual mean

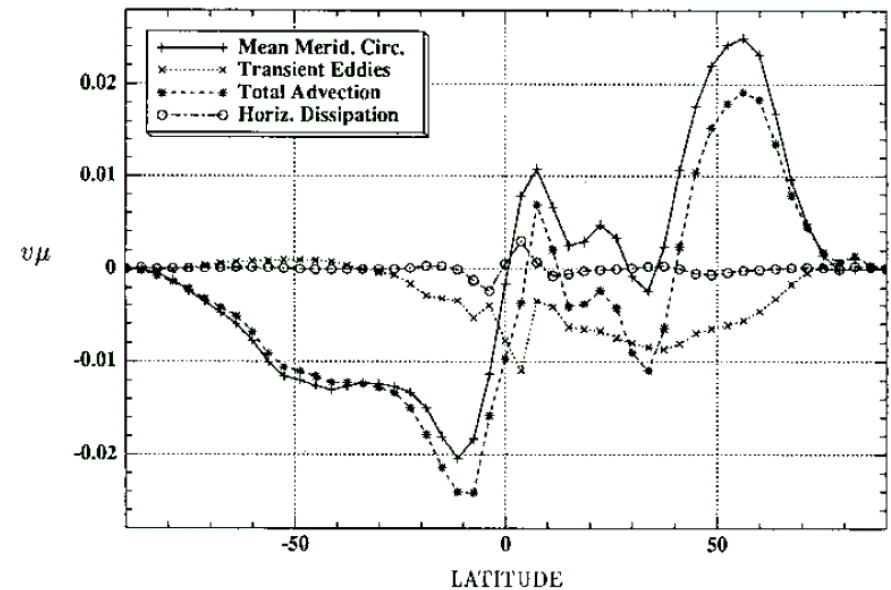


Momentum transport

N winter solstice



Hourdin et al. 1995



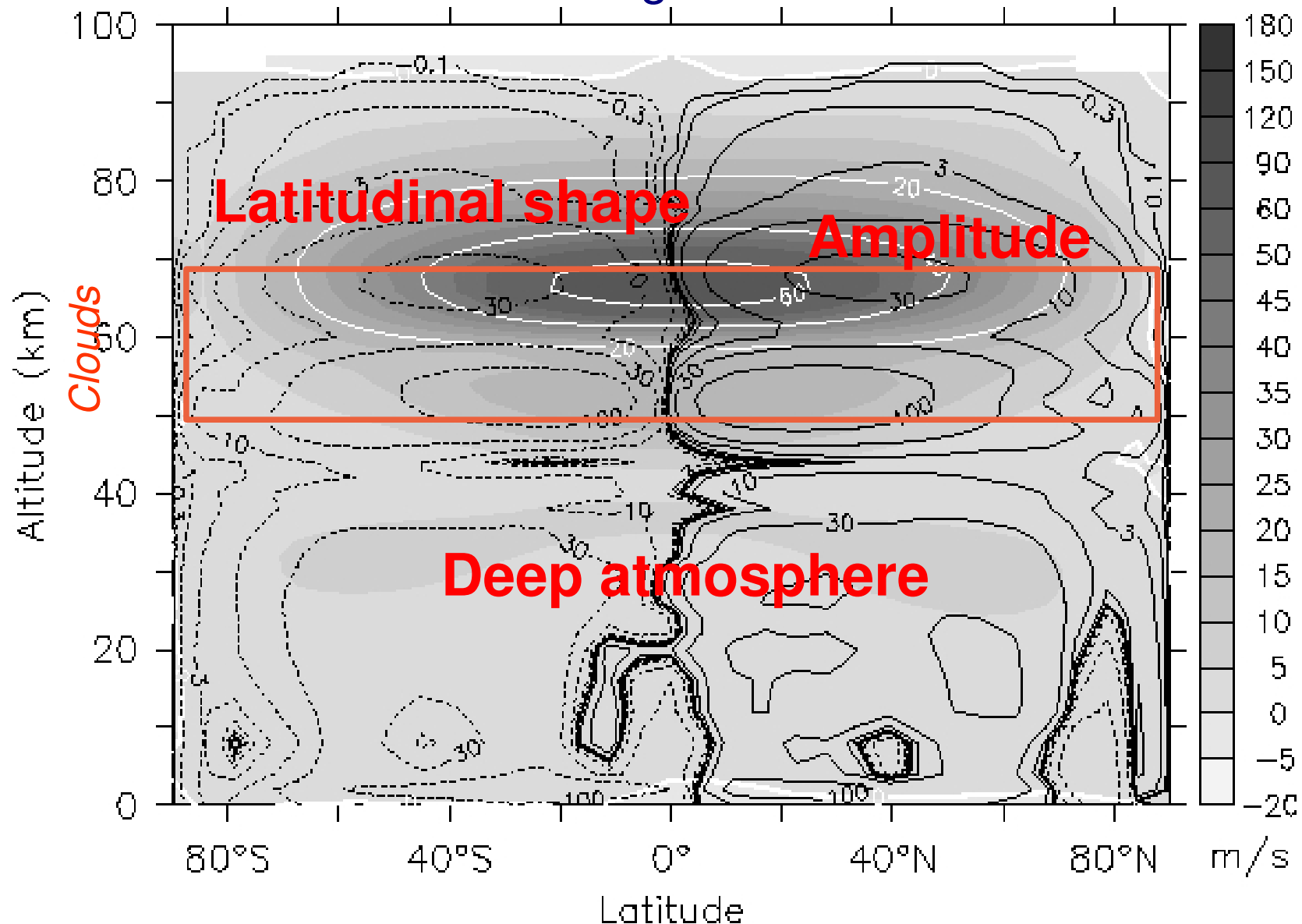
LMD VENUS GCM

- Three-dimensional: 48x32x50 (0~95 km)
- Vertical coordinates: hybrid (sigma/pressure)
- Dynamical core, transport of tracers
- Specific physics:
 - radiative transfer: Net Exchange Rates matrix
 - parameterizations (sub-grid processes, boundary layer, convection, turbulence)
 - topography
 - no clouds microphysics
- No photochemistry

Lebonnois et al., JGR, 2010

Venus Superrotation

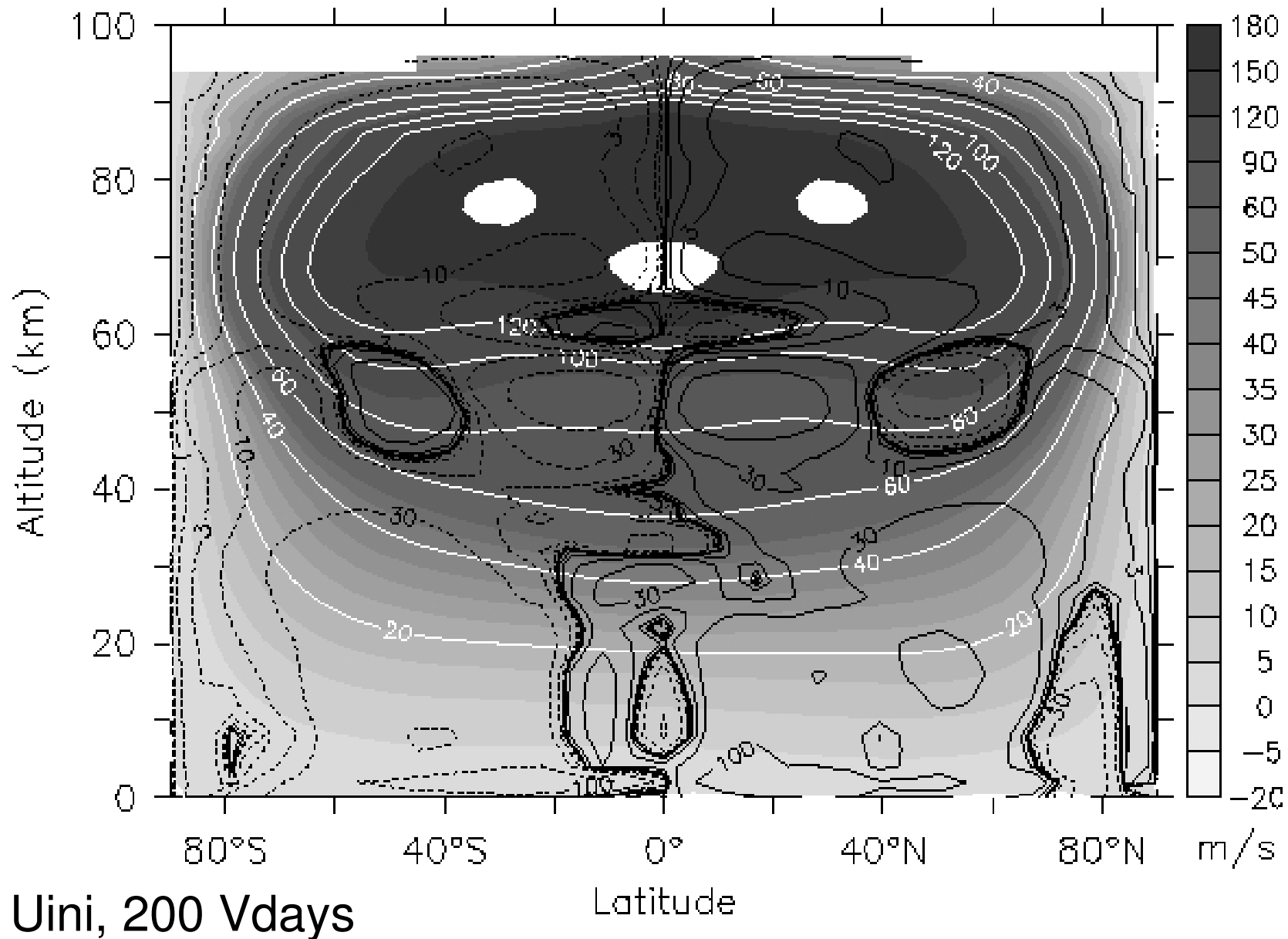
Starting from rest



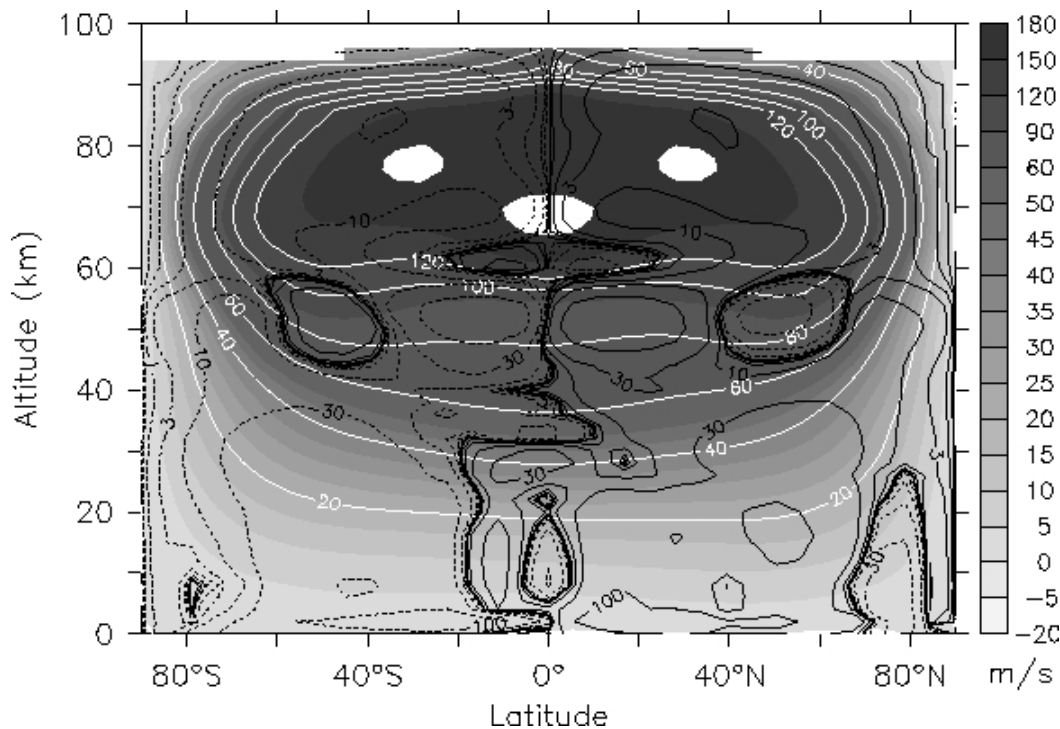
Mean zonal wind and stream function after 350 Vdays
(Topography, diurnal cycle)

Initial conditions

Starting from a zonal wind profile close to observations



Role of the diurnal cycle



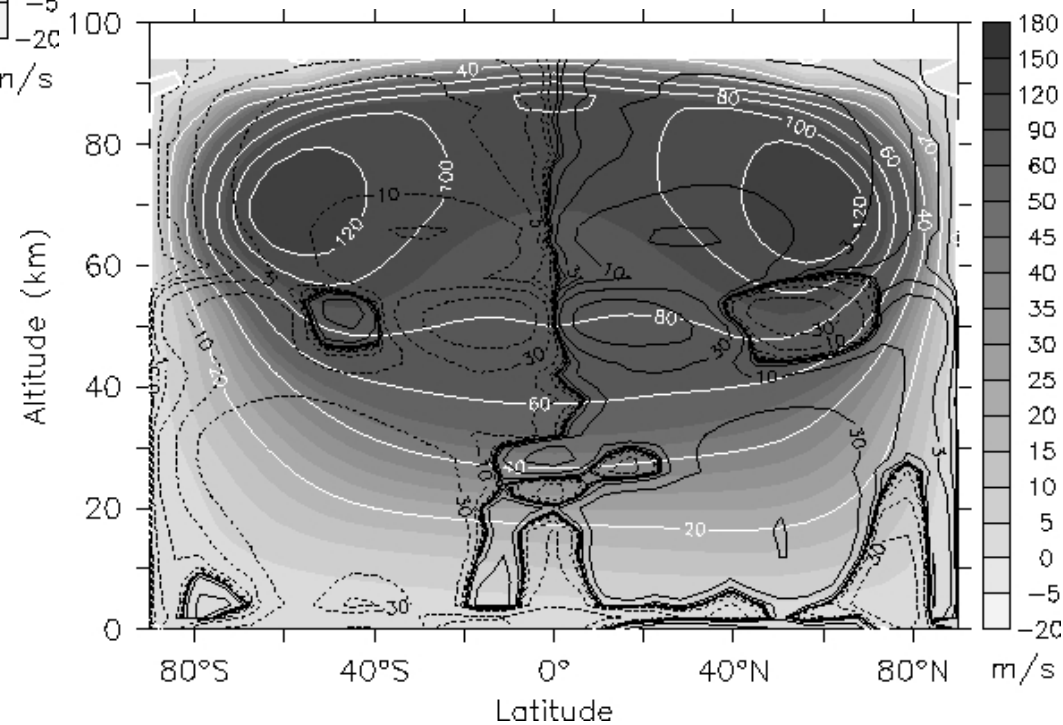
Uini With diurnal cycle

Influence of thermal tides in
angular momentum transport:

downward transport in the
equatorial 64-90 km region.

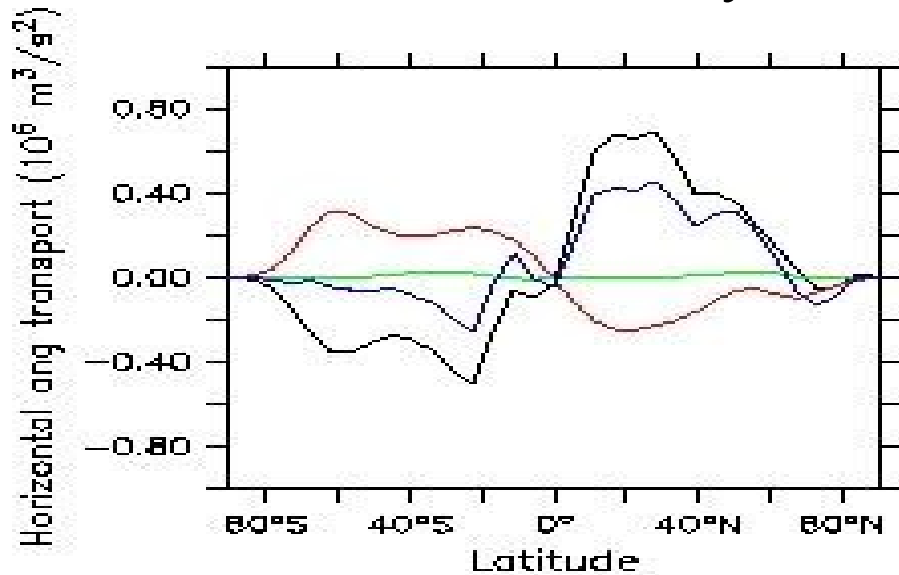
Undc

Without diurnal cycle

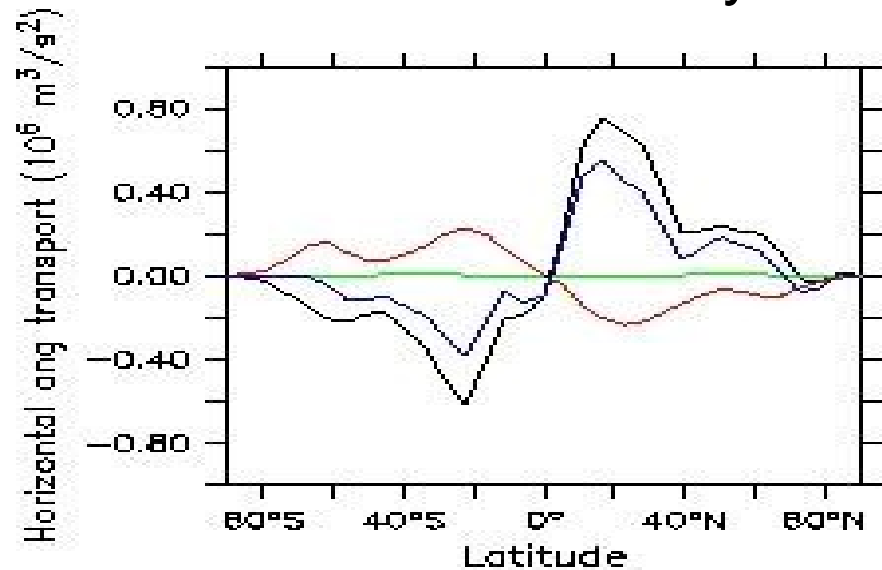


Angular momentum transport

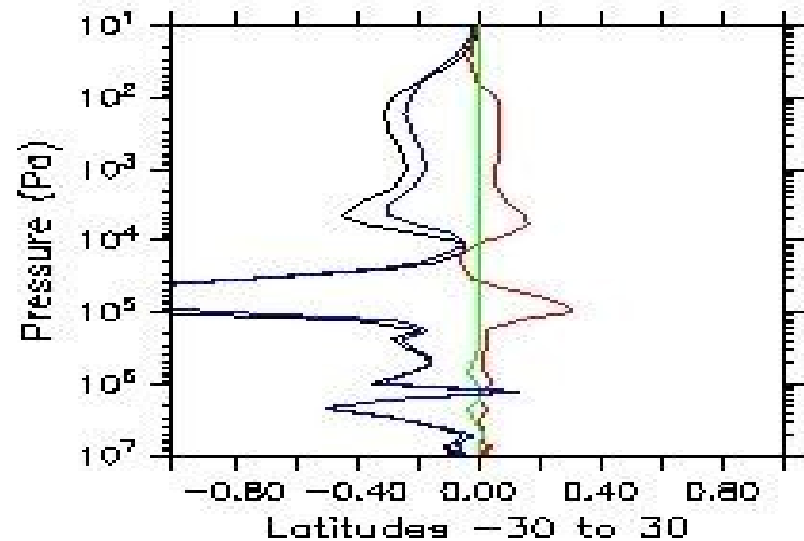
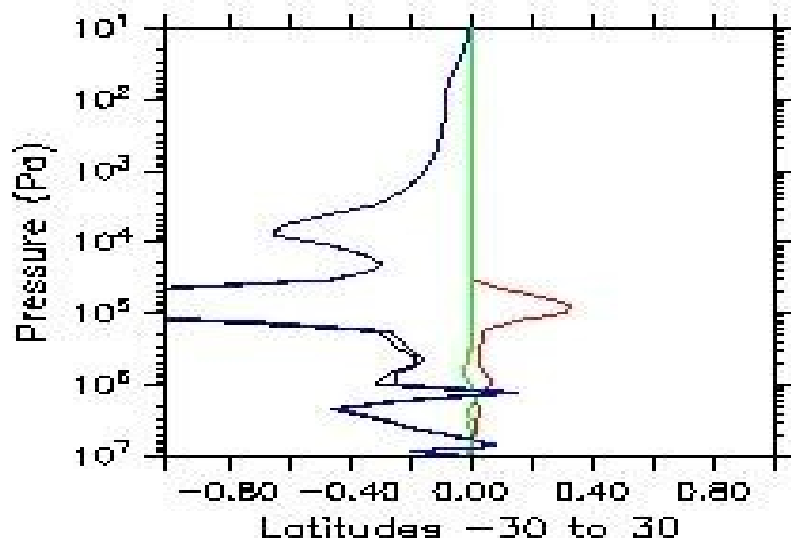
Without diurnal cycle



With diurnal cycle



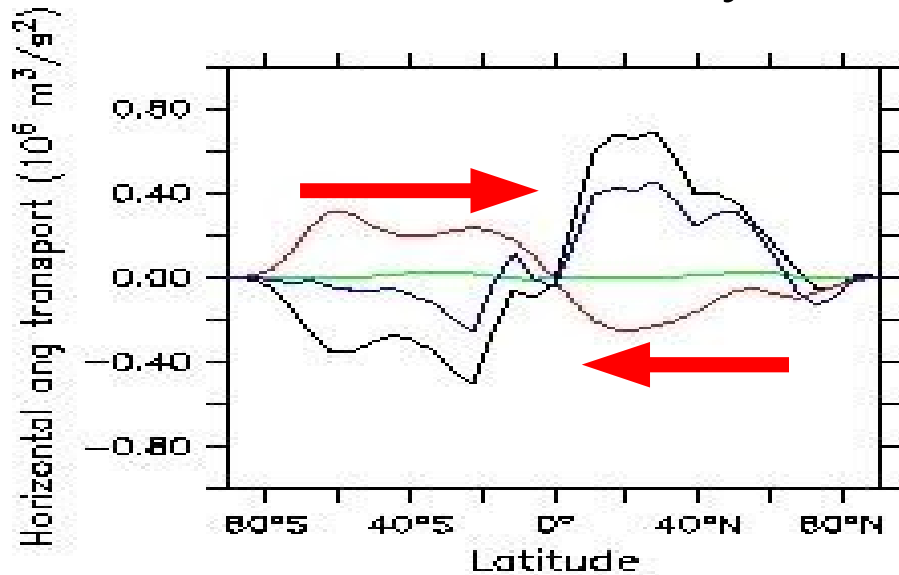
Vertical transport between 30°N and 30°S



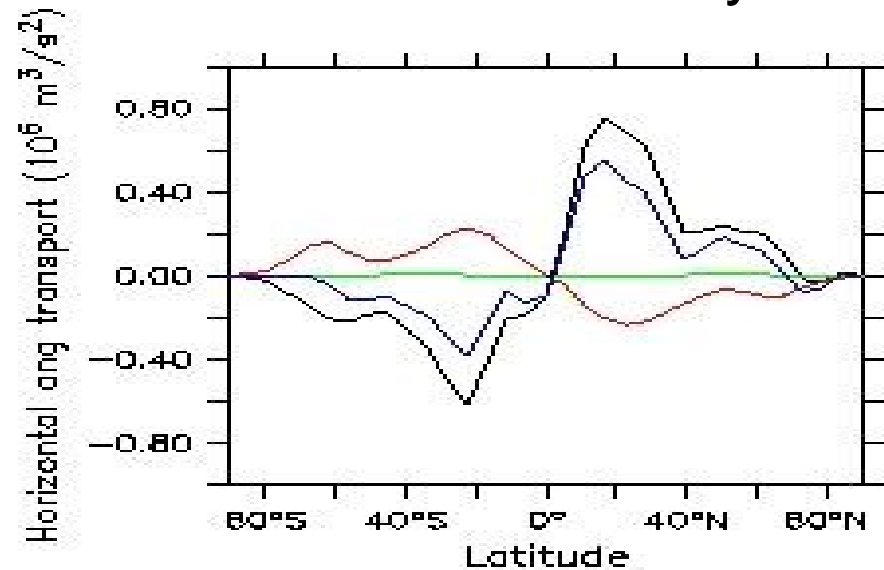
Angular momentum transport

Role of waves

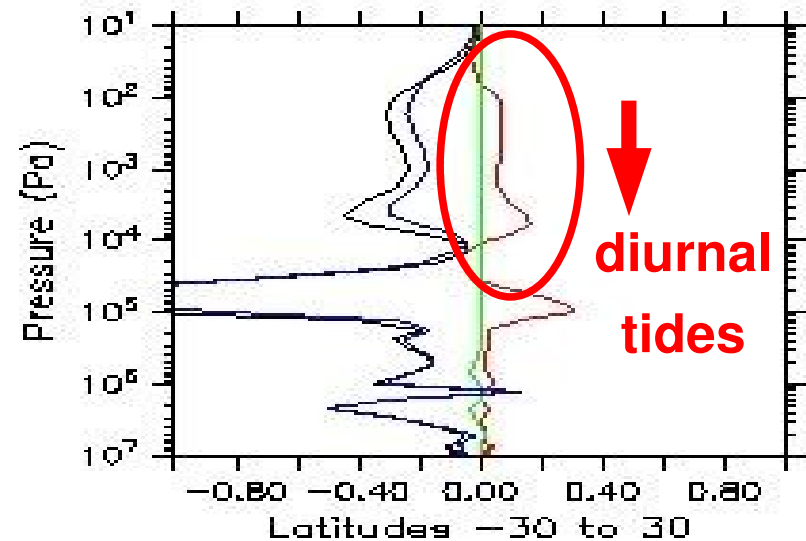
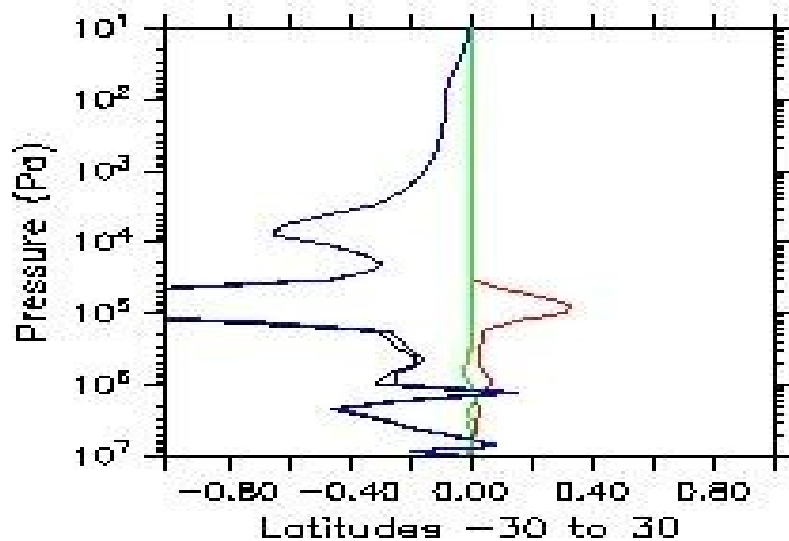
Without diurnal cycle



With diurnal cycle



Vertical transport between 30°N and 30°S



Discussion

- The superrotation problem is a difficult and sensitive one: many GCM have tried and failed to produce superrotation either for Venus or Titan, some succeed. Why ?
- Meridional circulation resulting from slow rotation.
- Titan vs Venus: the influence of seasonal variations.
- Non-axisymmetric angular momentum transport
 - Venus:** vertical transport in the equatorial region generated by thermal tides.
 - Titan:** unstable jet generating horizontal transport by waves. No role from thermal tides (radiative timescales).
- Venus: the question of the deep atmosphere is still pending...
What is missing ? Gravity waves forcing ?
Relation with Titan's model ?